

Full waveform modeling

Introduction

Geophysical imaging: to do what?

Generalities on Inverse Problems

Seismic data

A first glance at seismic inversion methods

Full waveform modeling

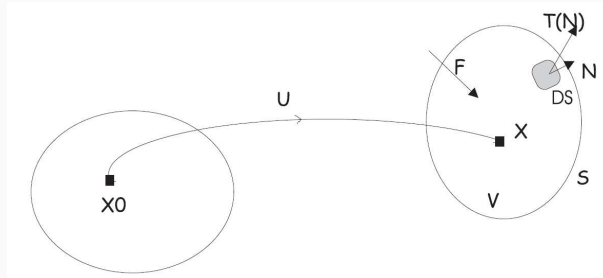
Building the wave equation

Heterogeneity, anisotropy and attenuation

Full waveform inversion

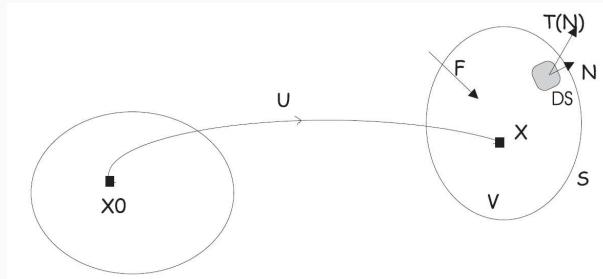
Displacement of a point

- reference system fixed in space and time
- a small element material is considered
- a point P at $\mathbf{x}_0 = (x_0, y_0, z_0)^t$ is considered at time t_0
- the point moves to $\mathbf{x}$ at time t



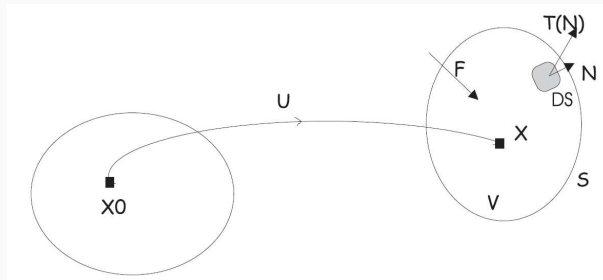
Displacement of a point

- displacement vector : $\mathbf{u} = \mathbf{x} - \mathbf{x}_0$
- velocity vector : $\frac{\partial \mathbf{u}}{\partial t}$
- acceleration vector $\frac{\partial^2 \mathbf{u}}{\partial t^2}$



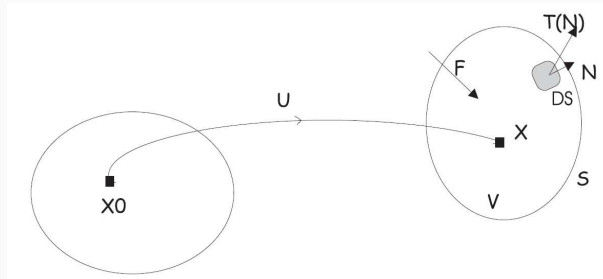
- we consider a close volume V , with mass conservation

$$\frac{dM(t)}{dt} = 0. \quad (8)$$



$$\iiint_V \rho \frac{\partial^2 \mathbf{u}}{\partial t^2} dV = \iiint_V \mathbf{F} dV + \iint_S \mathbf{T} dS. \quad (9)$$

- ρ : density
- $\mathbf{F}$: density of volumetric forces
- $\mathbf{T}$: surface forces

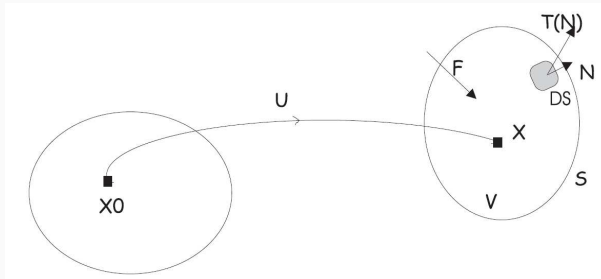


Simplification: surface forces $\rightarrow$ stress

- We can write

$$T_i = \sum_j \sigma_{ij} n_j, \quad (10)$$

- σ_{ij} : component ij of the stress tensor
- n_j : j component of the normal vector to S

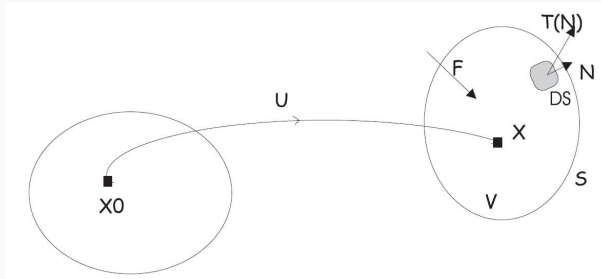


Simplification : divergence theorem

- divergence theorem

$$\iint_S \sigma_{ij} n_j dS = \iiint_V \sum_j \partial_j \sigma_{ij} dV \quad (11)$$

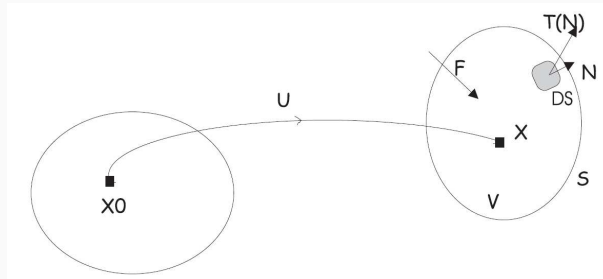
- ∂_j : spatial partial derivative along j



Second Newton law : integration on an elementary volume

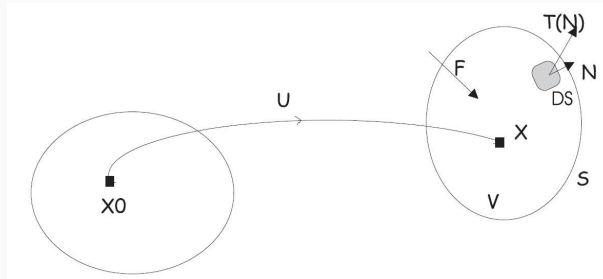
$$\rho \frac{\partial^2 u_i}{\partial t^2} = F_i + \sum_j \partial_j \sigma_{ij} \quad (12)$$

- i : component of the vector
- j : spatial direction



Second Newton law : expended version

$$\begin{aligned}\rho \frac{\partial^2 u_x}{\partial t^2} &= \frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} + F_x \\ \rho \frac{\partial^2 u_y}{\partial t^2} &= \frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} + F_y \\ \rho \frac{\partial^2 u_z}{\partial t^2} &= \frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} + F_z\end{aligned}\tag{13}$$



- linear part

$$\epsilon_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right) \quad (15)$$

→ linear relation between strain and displacement

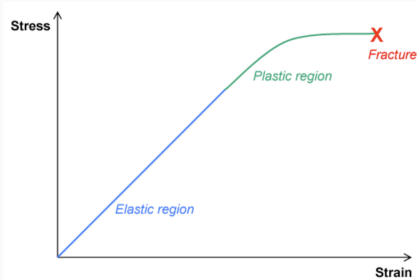


- rheology law between stress and strain

$$\sigma_{ij} = \sigma_{ij}^0 + C_{ijkl}\epsilon_{kl} + D_{ijklmn}\epsilon_{kl}\epsilon_{mn} + \mathcal{O}(\epsilon^3) \quad (16)$$

- σ_{ij}^0 : pre-stress
- C_{ijkl} et D_{ijklmn} : order 4 and 6 tensors
- linear part

$$\sigma_{ij} = \sigma_{ij}^0 + C_{ijkl}\epsilon_{kl} \quad (17)$$



- dynamic equation

$$\rho \frac{\partial^2 u_i}{\partial t^2} = F_i + \sum_j \partial_j \sigma_{ij} \quad (18)$$

- strain/displacement relation

$$\epsilon_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right) \quad (19)$$

- rheology relation

$$\sigma_{ij} = \sigma_{ij}^0 + C_{ijkl} \epsilon_{kl} \quad (20)$$

Voigt notations

$$11 \rightarrow 1, 22 \rightarrow 2, 33 \rightarrow 3, 23 \rightarrow 4, 13 \rightarrow 5, 12 \rightarrow 6. \quad (21)$$

This makes

$$\sigma = [\sigma_{11} \ \sigma_{22} \ \sigma_{33} \ \sigma_{23} \ \sigma_{13} \ \sigma_{12}]^T = [\sigma_1 \ \sigma_2 \ \sigma_3 \ \sigma_4 \ \sigma_5 \ \sigma_6]^T, \quad (22)$$

$$\epsilon = [\epsilon_{11} \ \epsilon_{22} \ \epsilon_{33} \ 2\epsilon_{23} \ 2\epsilon_{13} \ 2\epsilon_{12}]^T = [\epsilon_1 \ \epsilon_2 \ \epsilon_3 \ \epsilon_4 \ \epsilon_5 \ \epsilon_6]^T. \quad (23)$$

We introduce the matrix operator D as

$$D = \begin{pmatrix} \partial_{x_1} & 0 & 0 & 0 & \partial_{x_2} & \partial_{x_3} \\ 0 & \partial_{x_2} & 0 & \partial_{x_3} & 0 & \partial_{x_1} \\ 0 & 0 & \partial_{x_3} & \partial_{x_2} & \partial_{x_1} & 0 \end{pmatrix}. \quad (24)$$

Gathering equations we obtain

$$\begin{aligned}\rho \frac{\partial^2 u}{\partial t^2} &= D\sigma + F \\ \sigma &= \sigma^0 + CD^T u.\end{aligned}\tag{25}$$

- In elastic isotropic media, C_{ijkl} depends on 2 independent coefficients (λ and μ , or E and ν), giving

$$\begin{aligned}\sigma_{xx} &= (\lambda + 2\mu) \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{xx0} \\ \sigma_{yy} &= \lambda \frac{\partial u_x}{\partial x} + (\lambda + 2\mu) \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{yy0} \\ \sigma_{zz} &= \lambda \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + (\lambda + 2\mu) \frac{\partial u_z}{\partial z} + \sigma_{zz0} \\ \sigma_{xy} &= \mu \left(\frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x} \right) + \sigma_{xy0} \\ \sigma_{yz} &= \mu \left(\frac{\partial u_y}{\partial z} + \frac{\partial u_z}{\partial y} \right) + \sigma_{yz0} \\ \sigma_{xz} &= \mu \left(\frac{\partial u_x}{\partial z} + \frac{\partial u_z}{\partial x} \right) + \sigma_{xz0}\end{aligned}\tag{26}$$

- we combine the sets of equations

$$\begin{aligned}\rho \frac{\partial^2 u_x}{\partial t^2} &= \frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} + F_x \\ \rho \frac{\partial^2 u_y}{\partial t^2} &= \frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} + F_y \\ \rho \frac{\partial^2 u_z}{\partial t^2} &= \frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} + F_z \\ \sigma_{xx} &= (\lambda + 2\mu) \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{xx0} \\ \sigma_{yy} &= \lambda \frac{\partial u_x}{\partial x} + (\lambda + 2\mu) \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{yy0} \\ \sigma_{zz} &= \lambda \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + (\lambda + 2\mu) \frac{\partial u_z}{\partial z} + \sigma_{zz0} \\ \sigma_{xy} &= \mu \left(\frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x} \right) + \sigma_{xy0} \\ \sigma_{yz} &= \mu \left(\frac{\partial u_y}{\partial z} + \frac{\partial u_z}{\partial y} \right) + \sigma_{yz0} \\ \sigma_{xz} &= \mu \left(\frac{\partial u_x}{\partial z} + \frac{\partial u_z}{\partial x} \right) + \sigma_{xz0}\end{aligned}\tag{27}$$

- we combine the sets of equations
- PDE with 9 unknowns : 3 displacement + 3 stresses

$$\begin{aligned}\rho \frac{\partial^2 u_x}{\partial t^2} &= \frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} + F_x \\ \rho \frac{\partial^2 u_y}{\partial t^2} &= \frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} + F_y \\ \rho \frac{\partial^2 u_z}{\partial t^2} &= \frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} + F_z \\ \sigma_{xx} &= (\lambda + 2\mu) \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{xx0} \\ \sigma_{yy} &= \lambda \frac{\partial u_x}{\partial x} + (\lambda + 2\mu) \frac{\partial u_y}{\partial y} + \lambda \frac{\partial u_z}{\partial z} + \sigma_{yy0} \\ \sigma_{zz} &= \lambda \frac{\partial u_x}{\partial x} + \lambda \frac{\partial u_y}{\partial y} + (\lambda + 2\mu) \frac{\partial u_z}{\partial z} + \sigma_{zz0} \\ \sigma_{xy} &= \mu \left(\frac{\partial u_x}{\partial y} + \frac{\partial u_y}{\partial x} \right) + \sigma_{xy0} \\ \sigma_{yz} &= \mu \left(\frac{\partial u_y}{\partial z} + \frac{\partial u_z}{\partial y} \right) + \sigma_{yz0} \\ \sigma_{xz} &= \mu \left(\frac{\partial u_x}{\partial z} + \frac{\partial u_z}{\partial x} \right) + \sigma_{xz0}\end{aligned}\tag{27}$$

- we replace
 $v_i = \partial_t u_i$

$$\begin{aligned}\frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\ \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\ \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z\end{aligned}$$

(28)

- we replace
 $v_i = \partial_t u_i$
- we derive over
time the rheology
relation

$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\
 \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z \\
 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \frac{\partial \sigma_{yy}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{yy0}}{\partial t} \\
 \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + (\lambda + 2\mu) \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t} \\
 \frac{\partial \sigma_{xy}}{\partial t} &= \mu \left(\frac{\partial v_x}{\partial y} + \frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\
 \frac{\partial \sigma_{yz}}{\partial t} &= \mu \left(\frac{\partial v_y}{\partial z} + \frac{\partial v_z}{\partial y} \right) + \frac{\partial \sigma_{yz0}}{\partial t} \\
 \frac{\partial \sigma_{xz}}{\partial t} &= \mu \left(\frac{\partial v_x}{\partial z} + \frac{\partial v_z}{\partial x} \right) + \frac{\partial \sigma_{xz0}}{\partial t}
 \end{aligned} \tag{28}$$

Isotropic medium example

- we replace
 $v_i = \partial_t u_i$
- we derive over
time the rheology
relation

we obtain an homogeneous system of order 1 in time

$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xy}}{\partial y} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yy}}{\partial y} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\
 \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial y} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z \\
 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \frac{\partial \sigma_{yy}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{yy0}}{\partial t} \\
 \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + (\lambda + 2\mu) \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t} \\
 \frac{\partial \sigma_{xy}}{\partial t} &= \mu \left(\frac{\partial v_x}{\partial y} + \frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\
 \frac{\partial \sigma_{yz}}{\partial t} &= \mu \left(\frac{\partial v_y}{\partial z} + \frac{\partial v_z}{\partial y} \right) + \frac{\partial \sigma_{yz0}}{\partial t} \\
 \frac{\partial \sigma_{xz}}{\partial t} &= \mu \left(\frac{\partial v_x}{\partial z} + \frac{\partial v_z}{\partial x} \right) + \frac{\partial \sigma_{xz0}}{\partial t}
 \end{aligned} \tag{28}$$

- in acoustic medium, shear effects vanish

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$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \cancel{\frac{\partial \sigma_{xy}}{\partial y}} + \cancel{\frac{\partial \sigma_{xz}}{\partial z}} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\cancel{\frac{\partial \sigma_{xy}}{\partial x}} + \frac{\partial \sigma_{yy}}{\partial y} + \cancel{\frac{\partial \sigma_{yz}}{\partial z}} \right) + f_y \\
 \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\cancel{\frac{\partial \sigma_{xz}}{\partial x}} + \cancel{\frac{\partial \sigma_{yz}}{\partial y}} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z \\
 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + \cancel{2\mu}) \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \frac{\partial \sigma_{yy}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + \cancel{2\mu}) \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{yy0}}{\partial t} \\
 \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + (\lambda + \cancel{2\mu}) \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{xy}}{\partial t}} &= \cancel{\mu} \left(\frac{\partial v_x}{\partial y} + \frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{yz}}{\partial t}} &= \cancel{\mu} \left(\frac{\partial v_y}{\partial z} + \frac{\partial v_z}{\partial y} \right) + \frac{\partial \sigma_{yz0}}{\partial t} \\
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 \end{aligned} \tag{29}$$

- in acoustic medium, shear effects vanish

$$\begin{aligned}\frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{xx}}{\partial x} + f_x \\ \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{yy}}{\partial y} + f_y \\ \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{zz}}{\partial z} + f_z \\ \frac{\partial \sigma_{xx}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\ \frac{\partial \sigma_{yy}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{yy0}}{\partial t} \\ \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t}\end{aligned}\tag{30}$$

- in acoustic medium, shear effects vanish

$$\begin{aligned}\frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \frac{\partial P}{\partial x} + f_x \\ \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \frac{\partial P}{\partial y} + f_y \\ \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \frac{\partial P}{\partial z} + f_z \\ \frac{\partial P}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_y}{\partial y} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial P_0}{\partial t}\end{aligned}\quad (31)$$

hypothesis

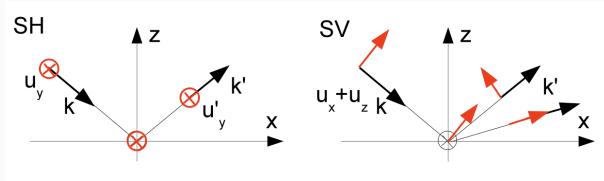
- invariant properties along one direction
- source and receivers invariant along one direction : line source and receivers

- invariance along y
(for example)

$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \cancel{\frac{\partial \sigma_{xy}}{\partial y}} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \cancel{\frac{\partial \sigma_{yy}}{\partial y}} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\
 \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \cancel{\frac{\partial \sigma_{yz}}{\partial y}} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z \\
 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{xy}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \cancel{\frac{\partial v_y}{\partial y}} + \lambda \frac{\partial v_z}{\partial z} + \cancel{\frac{\partial \sigma_{xy0}}{\partial t}} \\
 \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + (\lambda + 2\mu) \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t} \\
 \frac{\partial \sigma_{xy}}{\partial t} &= \mu \left(\cancel{\frac{\partial v_x}{\partial y}} + \frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\
 \frac{\partial \sigma_{yz}}{\partial t} &= \mu \left(\frac{\partial v_y}{\partial z} + \cancel{\frac{\partial v_z}{\partial y}} \right) + \frac{\partial \sigma_{yz0}}{\partial t} \\
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 \end{aligned} \tag{32}$$

- invariance along y
(for example)
- 2 independent systems appear

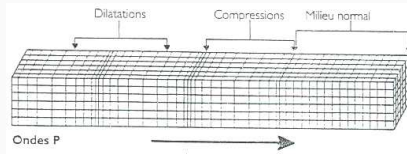
$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \cancel{\frac{\partial \sigma_{xy}}{\partial y}} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \cancel{\frac{\partial \sigma_{yy}}{\partial y}} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\
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 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{xy}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \cancel{\frac{\partial v_y}{\partial y}} + \lambda \frac{\partial v_z}{\partial z} + \cancel{\frac{\partial \sigma_{xy0}}{\partial t}} \\
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 \end{aligned} \tag{32}$$



2D isotropic medium : P-SV system

- P-SV system
 - P waves
 - S waves polarized in plane xz

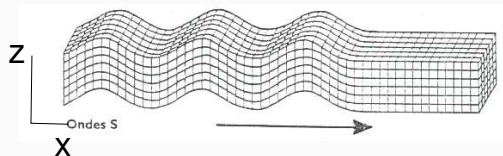
$$\begin{aligned}\frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xz}}{\partial z} \right) + f_x \\ \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{zz}}{\partial z} \right) + f_z \\ \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{xx0}}{\partial t} \\ \frac{\partial \sigma_{zz}}{\partial t} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \frac{\partial v_z}{\partial z} + \frac{\partial \sigma_{zz0}}{\partial t} \\ \frac{\partial \sigma_{xz}}{\partial t} &= \mu \left(\frac{\partial v_x}{\partial z} + \frac{\partial v_z}{\partial x} \right) + \frac{\partial \sigma_{xz0}}{\partial t}.\end{aligned}\quad (33)$$



2D isotropic medium : P-SV system

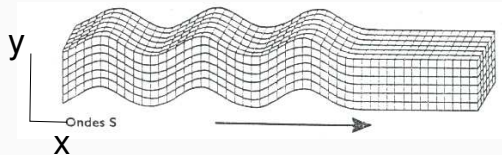
- P-SV system
 - P waves
 - S waves polarized in plane xz

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- SH system
 - S waves only
in plane xy

$$\begin{aligned}\frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \frac{\partial \sigma_{yz}}{\partial z} \right) + f_y \\ \frac{\partial \sigma_{xy}}{\partial t} &= \mu \left(\frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\ \frac{\partial \sigma_{yz}}{\partial t} &= \mu \left(\frac{\partial v_y}{\partial z} \right) + \frac{\partial \sigma_{yz0}}{\partial t}.\end{aligned}\tag{34}$$



Hypothesis

- invariant properties in 2 directions
- invariant source and receiver in 2 directions : plane wave

- y and z directions for invariance (for example)

$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \cancel{\frac{\partial \sigma_{xy}}{\partial y}} + \cancel{\frac{\partial \sigma_{xz}}{\partial z}} \right) + f_x \\
 \frac{\partial v_y}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xy}}{\partial x} + \cancel{\frac{\partial \sigma_{yy}}{\partial y}} + \cancel{\frac{\partial \sigma_{yz}}{\partial z}} \right) + f_y \\
 \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xz}}{\partial x} + \cancel{\frac{\partial \sigma_{yz}}{\partial y}} + \cancel{\frac{\partial \sigma_{zz}}{\partial z}} \right) + f_z \\
 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + \lambda \cancel{\frac{\partial v_z}{\partial z}} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{yy}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \cancel{\frac{\partial v_y}{\partial y}} + \lambda \cancel{\frac{\partial v_z}{\partial z}} + \cancel{\frac{\partial \sigma_{yy0}}{\partial t}} \\
 \cancel{\frac{\partial \sigma_{zz}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + (\lambda + 2\mu) \cancel{\frac{\partial v_z}{\partial z}} + \cancel{\frac{\partial \sigma_{zz0}}{\partial t}} \\
 \frac{\partial \sigma_{xy}}{\partial t} &= \mu \left(\cancel{\frac{\partial v_x}{\partial y}} + \frac{\partial v_y}{\partial x} \right) + \frac{\partial \sigma_{xy0}}{\partial t} \\
 \frac{\partial \sigma_{yz}}{\partial t} &= \mu \left(\cancel{\frac{\partial v_y}{\partial z}} + \cancel{\frac{\partial v_z}{\partial y}} \right) + \frac{\partial \sigma_{yz0}}{\partial t} \\
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 \end{aligned} \tag{35}$$

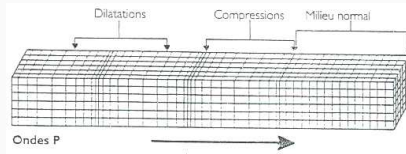
- y and z directions for invariance (for exemple)
- 3 independent systems appear

$$\begin{aligned}
 \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \left(\frac{\partial \sigma_{xx}}{\partial x} + \cancel{\frac{\partial \sigma_{xy}}{\partial y}} + \cancel{\frac{\partial \sigma_{xz}}{\partial z}} \right) + f_x \\
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 \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + \lambda \cancel{\frac{\partial v_z}{\partial z}} + \frac{\partial \sigma_{xx0}}{\partial t} \\
 \cancel{\frac{\partial \sigma_{yy}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + (\lambda + 2\mu) \cancel{\frac{\partial v_y}{\partial y}} + \lambda \cancel{\frac{\partial v_z}{\partial z}} + \cancel{\frac{\partial \sigma_{yy0}}{\partial t}} \\
 \cancel{\frac{\partial \sigma_{zz}}{\partial t}} &= \lambda \frac{\partial v_x}{\partial x} + \lambda \cancel{\frac{\partial v_y}{\partial y}} + (\lambda + 2\mu) \cancel{\frac{\partial v_z}{\partial z}} + \cancel{\frac{\partial \sigma_{zz0}}{\partial t}} \\
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 \end{aligned} \tag{35}$$

1D isotropic medium: P wave

- P wave along x direction

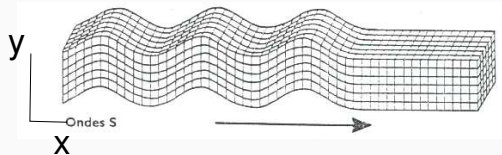
$$\begin{aligned}\frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{xx}}{\partial x} + f_x \\ \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \frac{\partial \sigma_{xx0}}{\partial t}\end{aligned}\quad (36)$$



1D isotropic medium: S wave (1)

- S wave along direction x , polarized in plane xy

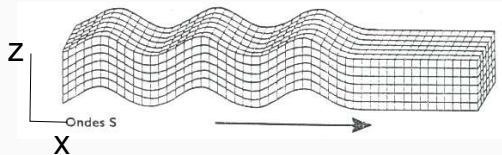
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1D isotropic medium: S wave (2)

- S wave along direction x , polarized in plane xz

$$\begin{aligned}\frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{xz}}{\partial x} + f_y \\ \frac{\partial \sigma_{xz}}{\partial t} &= \mu \frac{\partial v_z}{\partial x} + \frac{\partial \sigma_{xz0}}{\partial t}\end{aligned}\quad (38)$$



Which velocities for these waves ?

We can write the wave equation as a compact matrix-vector formulation

$$\frac{\partial \mathbf{U}}{\partial t} = A \frac{\partial \mathbf{U}}{\partial x} + B \frac{\partial \mathbf{U}}{\partial y} + C \frac{\partial \mathbf{U}}{\partial z} + \frac{\partial \mathbf{U}_0}{\partial t}, \quad (39)$$

where $\mathbf{U} = (v_x, v_y, v_z, \sigma_{xx}, \sigma_{yy}, \sigma_{zz}, \sigma_{xy}, \sigma_{yz}, \sigma_{xz})^t$.

Which velocities for these waves ?

Example in 1D

$$\frac{\partial \mathbf{U}}{\partial t} = A \frac{\partial \mathbf{U}}{\partial x} + \frac{\partial \mathbf{U}_0}{\partial t}, \quad (40)$$

- P wave

$$\begin{aligned} \frac{\partial v_x}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{xx}}{\partial x} + f_x \\ \frac{\partial \sigma_{xx}}{\partial t} &= (\lambda + 2\mu) \frac{\partial v_x}{\partial x} + \frac{\partial \sigma_{xx0}}{\partial t} \end{aligned}$$

- S wave (in plane xz)

$$\begin{aligned} \frac{\partial v_z}{\partial t} &= \frac{1}{\rho} \frac{\partial \sigma_{xz}}{\partial x} + f_y \\ \frac{\partial \sigma_{xz}}{\partial t} &= \mu \frac{\partial v_z}{\partial x} + \frac{\partial \sigma_{xz0}}{\partial t} \end{aligned}$$

Which velocities for these waves ?

- For P -wave system we have

$$A = \begin{pmatrix} 0 & \frac{1}{\rho} \\ \lambda + 2\mu & 0 \end{pmatrix} \quad (41)$$

therefore the eigenvalues are solution of

$$X^2 - \frac{\lambda + 2\mu}{\rho} = 0 \quad (42)$$

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$$I = \pm \sqrt{\frac{\lambda + 2\mu}{\rho}} \rightarrow v_P = \sqrt{\frac{\lambda + 2\mu}{\rho}} \quad (43)$$

Which velocities for these waves ?

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- S wave (in plane xz)

$$I = \pm \sqrt{\frac{\mu}{\rho}} \rightarrow v_S = \sqrt{\frac{\mu}{\rho}} \quad (44)$$

Introduction

Geophysical imaging: to do what?

Generalities on Inverse Problems

Seismic data

A first glance at seismic inversion methods

Full waveform modeling

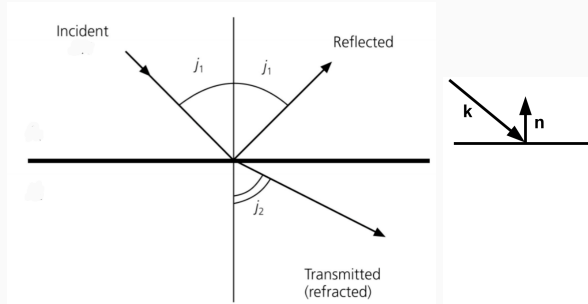
Building the wave equation

Heterogeneity, anisotropy and attenuation

Full waveform inversion

Reflection and refraction : Snell-Descartes law (1)

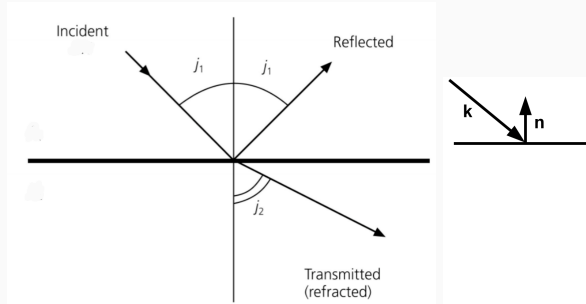
As for light...



$$\frac{\sin j_1}{v_1} = \frac{\sin j_2}{v_2} \quad j_1 = j'_1 \quad (45)$$

Reflection and refraction : Snell-Descartes law (1)

As for light...

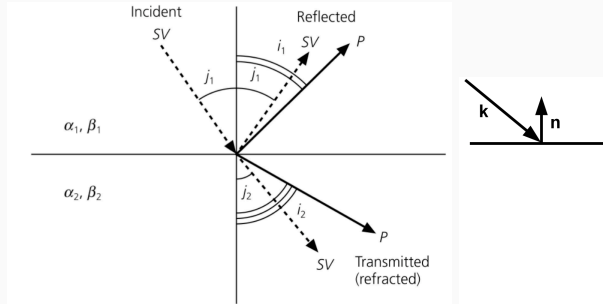


$$\frac{\sin j_1}{v_1} = \frac{\sin j_2}{v_2} \quad j_1 = j_1' \quad (45)$$

- this setup is valid for single wave type
 - P waves in acoustic media
 - SH waves (no coupling with P waves)

Reflection and refraction : Snell-Descartes law (2)

As for light...

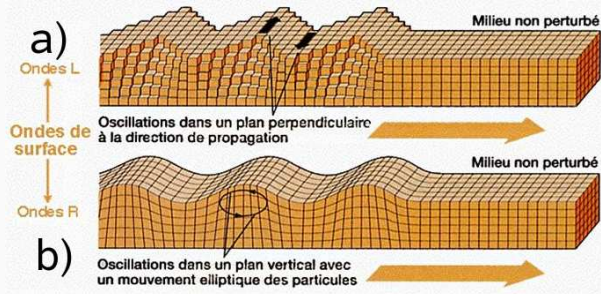


$$\frac{\sin j_1}{V_{S1}} = \frac{\sin j_2}{V_{S2}} = \frac{\sin i_1}{V_{P1}} = \frac{\sin i_2}{V_{P2}} \quad j_1 = j'_1 \quad (46)$$

- this setup is generic
 - P waves in solids
 - SV waves

- the free surface is a particular interface
 - on one side a solid : v_{P_S} , v_{S_S} and ρ_S
 - on the other side, air : $v_{P_{air}} < v_{P_S}$, $v_S = 0$ and $\rho_{air} \ll \rho_S$

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 - on the other side, air : $v_{P_{air}} < v_{P_S}$, $v_S = 0$ and $\rho_{air} \ll \rho_S$
→ we generally assume air as void



Rayleigh waves

- particles move in the P-SV plane
- result from interferences between P and SV waves

$$v_R < v_S < v_P \quad (47)$$

$$v_R^6 - 8v_S^2 v_R^4 + (24 - 16v_S^2/v_P^2)v_S^4 v_R^2 + 16(v_S^2/v_P^2 - 1)v_S^6 = 0 \quad (48)$$

where v_R is the Rayleigh wave velocity

Rayleigh waves

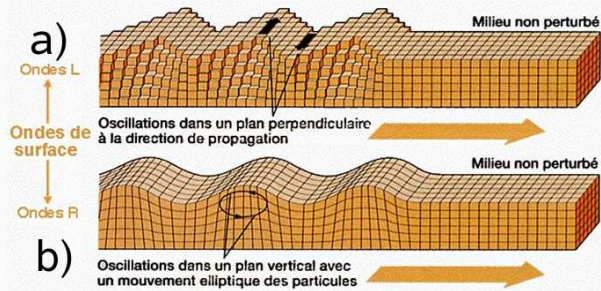
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where v_R is the Rayleigh wave velocity

for perfect solid media, $\nu \approx 0.25$ giving $v_R = 0.919v_S$

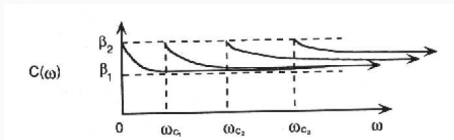
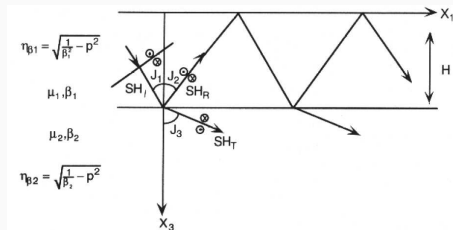


Love waves

- particle motion in the SH plane
- result from interferences between incident, reflected and refracted SH in an heterogeneous near surface
→ does not exist in homogeneous media
- we have

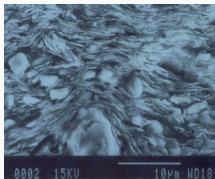
$$v_{S1} < v_L < v_{S2} \quad (49)$$

- giving dispersion in all cases

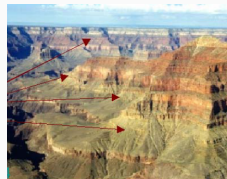


- Anisotropy : changing of the behavior depending on the direction
- In geologic media, wave anisotropy can have two origins : internal (mineral composition → static anisotropy) and external (layer structure for example → dynamic anisotropy)

Micro-scale anisotropy: mineral anisotropy



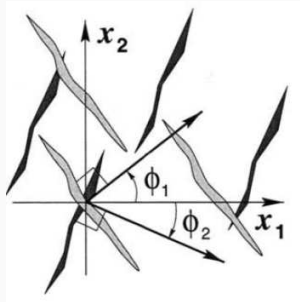
Macro-scale anisotropy: VTI/TTI anisotropy



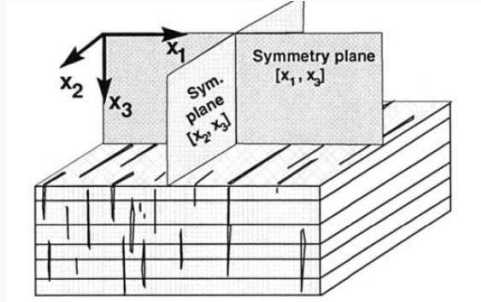
- in all cases, anisotropy is a scale problem : a medium is seen as anisotropic when it contains “small scale” heterogeneities or structures
→ “small scale” is related to wavelength $\lambda = V/f$ for wave propagation

Anisotropy

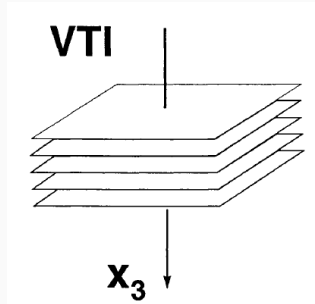
- In anisotropic media, the reologic relation is more complex
 - **triclinic** : 21 independent coefficients in C_{ijkl}
 - **monoclinic** : 13 independent coefficients in C_{ijkl}



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 - **triclinic** : 21 independent coefficients in C_{ijkl}
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 - **orthorombic** : 9 independent coefficients in C_{ijkl}

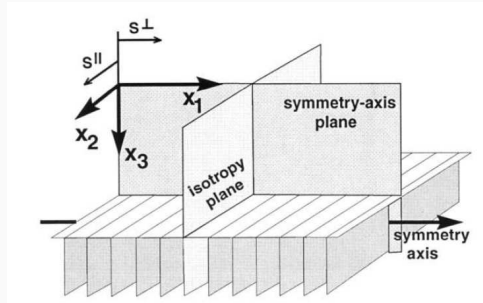


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Anisotropy

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Attenuation: intrinsic, extrinsic and apparent

- Quality factor : relate the energy loss per wavelength (effect is frequency-dependent)
- energy loss can be related to intrinsic (fluid motion in porous media, plastic deformation, anelastic reology...) or extrinsic (multi-scattering in heterogeneous media)

Shallow layer

Strong attenuation

Difficult to find equivalent visco-elastic model

Crust and Mantle

Visco-elastic model of anelasticity

Rheological laws (near-elastic & plastic deformation)

Frequency dependence of the Q factor
(Pertinence of the Nearly Constant Q - NCQ- model)

